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1 Incorporating diffuse radiation into a light use efficiency and evapotranspiration model:

- 2 an 11-year study in a high latitude deciduous forest
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8 Abstract:

9 The fraction of diffuse photosynthetic active radiation (PAR) reaching the land surface is one of the biophysical 10 factors regulating carbon and water exchange between ecosystems and the atmosphere. This is especially relevant for high latitude ecosystems, where cloudy days are prevalent. Without considering impacts of diffuse 11 PAR, traditional 'top-down' models of ecosystem gross primary productivity (GPP) and evapotranspiration (ET), 12 13 which use satellite remote sensing observations, tend to be biased towards clear sky conditions. Thus, this study incorporated a cloudiness index (CI), an index for the fraction of diffuse PAR, into a joint 'top-down' model that 14 15 uses the same set of biophysical constraints to simulate GPP and ET for a high latitude temperate deciduous 16 forest. To quantify the diffuse PAR effects, CI along with other environmental variables derived from an eleven-17 year eddy covariance data set were used to statistically explore the independent and joint effects of diffuse PAR on GPP, ET, incident light use efficiency (LUE), evaporative fraction (EF) and ecosystem water use efficiency 18 (WUE). The independent and joint effects of CI were compared from global sensitivity analysis of the 'top-down' 19 models. Results indicate that for independent effects, CI increased GPP, LUE, ET, EF and WUE, but analysis of 20 joint effects shows that as CI mainly interacted with the radiation intercepted in the canopy (PAR, net radiation 21 and leaf area index) to influence GPP, ET and WUE. Moreover, Ta and vapor pressure deficit played a major 22 role for the joint influence of CI on LUE and EF. We quantified that CI contributes 11.88%, 3.04% and 7.78% to 23 24 the total variation of GPP, ET and transpiration in the growing season from May to October, respectively. As the 25 influence of CI on GPP is larger than that on ET, this leads to an increase in WUE. Joint GPP and ET model results showed that when including CI, the root mean square errors (RMSE) of daily GPP decreased from 1.64 to 26 27 1.45 g·C·m⁻²·d⁻¹ (11.68% reduction) and ET from 15.79 to 14.50 W·m⁻² (8.16% reduction). Due to the interaction of diffuse PAR with plant canopies, the largest model improvements using CI for GPP and ET occurred during 28 29 the growing season and for the transpiration component, as suggested by comparisons to sap flow measurements. Furthermore, our study suggests a potential biophysical mechanism, not considered in other studies: due to the 30 31 increased longwave emission from clouds, surface temperature gets higher and closer to optimum, boosting GPP 32 and transpiration in the temperature-limited high latitude ecosystem.

Key words: diffuse PAR fraction; eddy covariance; gross primary production; evapotranspiration; 'top-down'
 models; light use efficiency model; Priestley–Taylor Jet Propulsion Laboratory evapotranspiration model

- 36 Table of abbreviations and symbols:
- 37 Latin alphabet
- CI: cloudiness index (dimensionless)
- **39** EF: evaporative fraction (dimensionless)
- 40 ET: evapotranspiration (mm·d⁻¹)
- 41 f_{APAR} : fraction of absorbed PAR (dimensionless)
- 42 f_{ci} : cloudiness index constraint (dimensionless)
- 43 f_{diff} : fraction of diffuse PAR (dimensionless)
- f_g : the green canopy fraction indicating the proportion of active canopy (dimensionless)
- 45 f_M : the plant moisture constraint (dimensionless)
- 46 f_{IPAR} : fraction of intercepted PAR (dimensionless)
- 47 f_{Ta} : the air temperature constraint reflecting the temperature limitation of photosynthesis (dimensionless)
- 48 f_{SWC} : the soil moisture constraint on photosynthesis (dimensionless)
- 49 f_{VPD} : the VPD constraint reflecting the stomatal response to the atmospheric water saturation deficit 50 (dimensionless)
- 51 G: Ground heat flux $(W \cdot m^{-2})$
- GPP: gross primary productivity $(g \cdot C \cdot m^{-2} \cdot d^{-1})$
- 53 k_{PAR} : the extinction coefficients for PAR (0.5, dimensionless)
- k_{Rn} : the extinction coefficients for Rn (0.6, dimensionless)
- **55** LAI: leaf area index $(m^2 \cdot m^{-2})$
- LUE: incident light use efficiency $(g \cdot C \cdot MJ^{-1})$
- 57 LW_{in} : incoming longwave radiation (W·m⁻²)
- LW_{out}: outgoing longwave radiation (W·m⁻²)
- NDVI: normalized difference vegetation index (dimensionless)
- PAR: photosynthetically active radiation (MJ·m⁻²·d⁻¹)
- PARc: PAR intercepted by the canopy $(MJ \cdot m^{-2} \cdot d^{-1})$
- 62 RH: the relative humidity (dimensionless)
- 63 Rn: Net radiation $(W \cdot m^{-2})$
- Rnc: Net radiation intercepted by the canopy $(W \cdot m^{-2})$
- Rns: Net radiation reaching to the soil $(W \cdot m^{-2})$
- **66** SWC: soil water content $(m^3 \cdot m^{-3})$
- 67 SW_{in}: incoming shortwave radiation (W·m⁻²)

- SZA: sun zenith angle (rad)
- Ta: air temperature (°C)
- 70 Ts: surface temperature (°C)
- To: optimal air temperature for vegetation growth (°C)
- VPD: vapor pressure deficit (hPa)
- WUE: ecosystem water use efficiency $(g \cdot C \cdot kg^{-1})$
- 74

75 Greek alphabet

- 76 α: PT coefficient, an empirical ratio of potential evapotranspiration to equilibrium potential
- 77 evapotranspiration (dimensionless)
- γ : the psychrometric constant (0.066 kPa·°C⁻¹)
- **79** Δ : the slope of saturation-to-vapor pressure curve (kPa·°C⁻¹)
- 80 ε: surface emissivity (dimensionless)
- 81 ϵ_{max} : maximum LUE (g·C·m⁻²·MJ⁻¹)
- 82 λ : latent heat of vaporization (kJ·kg⁻¹)
- 83 λ ET: latent heat flux of evapotranspiration (W·m⁻²)
- λ Ec: latent heat flux from transpiration (W·m⁻²)
- 85 $\lambda Ei:$ latent heat flux from evaporation of intercepted water (W·m⁻²)
- 86 λ Es: latent heat flux from evaporation of soil water (W·m⁻²)
- 87 σ : the Stefan-Boltzmann constant (5.670367×10⁻⁸ kg·s⁻³·K⁻⁴)
- 88 89
- 90 1. Introduction

91 Quantifying land surface water and carbon fluxes is of critical importance for ecosystem and water resources 92 management. The temporal dynamics of land surface carbon and water fluxes are controlled by the interplay of 93 various biophysical factors, e.g. climate forcing (solar radiation, water vapor and temperature), atmospheric 94 conditions (CO₂ concentration and nitrogen deposition) and biotic factors (leaf area index and plant functional 95 types) (Ciais et al., 2005; Dunn et al., 2007; Wu et al., 2016). Among these biophysical factors, the fraction of 96 diffuse photosynthetically active radiation (PAR), f_{diff} (the ratio between diffuse and total PAR), has been 97 highlighted to have strong implications for the global carbon cycle (Gu et al., 2003; Mercado et al., 2009). It 98 could increase the efficiency of photosynthesis, which has been referred to the diffuse fertilization effect 99 (Roderick et al., 2001; Kanniah et al., 2012). Further, predictions showed that, at the global scale, aerosols in the atmosphere would increase by 36% in 2100 (Heald et al., 2008). Aerosols influence cloud formation and 100 101 increase f_{diff} in the atmosphere (Schiermeier, 2006). This is especially important for high latitude ecosystems, 102 which are already exposed to a higher f_{diff} due to low solar height and high frequency of overcast and cloudy conditions. 103

104 With more uniform vertical distribution of incoming photosynthetic active radiation (PAR) under cloudy conditions, both observations and modeling studies have confirmed more active carbon assimilation rates (Gu et 105 al., 2002; Lloyd et al., 2002; Steiner and Chameides, 2005; Ibrom et al. 2006, Urban et al., 2012). However, the 106 gross primary productivity (GPP) enhancement depends on local environmental conditions and ecosystem types. 107 108 Healy et al. (1998) reported that increasing f_{diff} can increase the incident light use efficiency (LUE, defined as the 109 ratio between GPP and incoming PAR). This increases crop yield by as much as 50% for maize, soybean and peanuts. According to observations from 10 temperate forest flux sites in USA, Cheng et al. (2015) found that 110 $f_{\rm diff}$ explained up to 41% and 17% of seasonal variations in GPP in croplands and forests, respectively. In a 111 112 modeling study, Ibrom et al. (2006) found the uniform PAR distribution in the maritime Scottish climate with a ca. 20% higher f_{diff} lead to a 13-14% higher LUE compared to the continental climate in Germany in spruce 113 canopies. To identify the impacts of f_{diff} , the covariance of f_{diff} and other environmental factors (Kanniah et al., 114 2012) should also be taken into account. For instance, Williams et al. (2016) found that without considering the 115 116 covariance between f_{diff} and phenology, the GPP enhancement from f_{diff} is 260%, while by separating f_{diff} and phenology, the GPP enhancement induced by f_{diff} dropped to 22%. Apart from modeling studies at the global 117 scale (Mercado et al., 2009), few studies have focused on ecosystems in high latitude regions, which are 118 119 radiation and temperature limited (van Dijk et al., 2005; Lagergren et al., 2008). In these ecosystems, the influence of f_{diff} and its covariance with other environmental variables should be thoroughly quantified, because 120 the potential mechanisms influencing GPP and ET might be different from those of water-limited ecosystems. 121

Because photosynthesis and transpiration are closely linked via stomatal behaviors, f_{diff} is expected to also have 122 123 moderate impacts on land evapotranspiration (ET) and may eventually influence the global hydrological cycle 124 and the climate system (Knohl and Baldocchi, 2008; Davin and Seneviratne, 2012; Pedruzo-Bagazgoitia et al., 125 2017). For instance, the modeling results from the Community Land Model showed that higher f_{diff} during 1960– 1990 increased the latent heat flux of evapotranspiration (λ ET) in the tropics by 2.5 Wm⁻² (3% of mean) and 126 127 reduced global river runoff (Oliveira et al., 2011). By employing the COSMO-CLM2 regional climate model, Davin and Seneviratne (2012) identified f_{diff} could alter the seasonal evaporative fraction (EF, defined as the ratio 128 129 between λET and available energy, which is net radiation minus soil heat flux Rn-G) and a consistent fraction 130 (up to 3%) of the overall variability in European summer air temperature could be explained by $f_{\rm diff}$. With increasing f_{diff} , the magnitude of the ET increase due to f_{diff} has been shown to be smaller than that of GPP, 131 resulting in an increase in the ecosystem water use efficiency (WUE, defined as the ratio between GPP and ET) 132 (Knohl and Baldocchi, 2008; Oliveira et al., 2011). Similarly to GPP, the local environment can also alter the 133 134 responses of ecosystem ET, EF and WUE to fdiff. For instance, in temperature-limited ecosystems at high latitudes, incoming longwave radiation has been shown to be an important source of energy for snow and glacier 135 136 melting under cloudy conditions with high f_{diff} increasing surface temperature (Juszak & Pelliciotti, 2013). However, the impacts of higher longwave radiation on the energy budget and canopy temperature have not been 137 138 considered yet, despite their potentially important implications for vegetation activities. In general, compared to studies on evaluating impacts of fdiff on GPP and LUE, studies on the influence of fdiff on ET, EF and WUE are 139 140 limited. More studies are needed to quantify impacts and understand mechanisms linking f_{diff} to ET, EF and WUE. 141

142 Traditionally, models that incorporate satellite remotely sensed observations, e.g. vegetation indices, surface 143 temperature or albedo, to estimate GPP and ET, tend to be biased to clear sky conditions, due to lack of 144 representation of cloudy conditions. These remote sensing models estimating GPP and ET can be classified into 145 'top-down' and 'bottom-up' approaches (Houborg et al., 2009). 'Top-down' methods, e.g. CASA (Potter et al., 1993), the MODIS GPP and ET algorithms (Running et al., 2004; Mu et al., 2007) or the Priestley–Taylor Jet 146 Propulsion Laboratory (PT-JPL) ET model (Fisher et al., 2008; Garcia et al., 2013), are simpler and can be 147 directly driven with remote sensing variables. These models try to represent the ecological behavior of the 148 149 canopy as a whole, using effective variables and few parameters. 'Top-down' models generally estimate GPP and ET assuming that the maximum LUE (ε_{max}) of plant canopies and the maximum ET are constrained by 150 similar stress-constraints reflecting different environmental constraints (Leuning et al., 1995; Houborg and 151 Soegaard, 2004; Houborg et al., 2009; Garcia et al., 2013). 'Top-down' GPP approaches, have the advantage that 152 153 forcing variables, such as the fraction of absorbed PAR at the top of the canopy or the land surface temperature, can be routinely estimated from remote sensing data instead of using a detailed description of canopy profiles 154 and leaf energy budgets as in 'bottom-up' methods (Wang and Leuning, 1998; Ryu et al., 2011). Due to the 155 impacts on the overall ecosystem GPP and ET, 'top-down' approaches can also benefit from considering f_{diff} . For 156 instance, Yuan et al. (2014) found six of the seven LUE GPP models, which did not consider fdiff, significantly 157 158 underestimated GPP during cloudy days. Recent studies incorporated the sunlit and shaded leaf approach into the 159 MODIS LUE algorithm to improve satellite based GPP estimation (He et al., 2013; Zhou et al., 2015). Donohue et al. (2014) extended Roderick et al.'s (2001) LUE formulation to include the dependency of ε_{max} on f_{diff} and on 160 the light-saturated rate of photosynthesis at the top of the canopy, yielding a highly generic model that accurately 161 predicted GPP across Australia. Wang et al. (2015) added the information of cloudiness index (CI) to improve 162 MODIS LUE algorithm. However, for 'top-down' ET models e.g. the Priestley Taylor based PT-JPL ET model, 163 the effects of f_{diff} have not been investigated previously. Furthermore, it has not been investigated, whether or not 164 the same set of biophysical constraints can be used to down-regulate both GPP and ET. 165

The outputs of 'top-down' models can be evaluated against eddy covariance (EC) datasets, including carbon and 166 167 water fluxes between the land surface and the atmosphere at the ecosystem scale. Long-term eddy covariance 168 and micrometeorological observations are also important to assess the environmental controls of carbon and 169 water exchange (Baldocchi et al., 2003) using statistical approaches, e.g. path analysis (Bassow & Bazzaz, 1998; 170 Huxman et al., 2003; Wu et al., 2016). In this study, we used a an 11-year time series of EC observations, a joint 171 GPP and ET 'top-down' model and in-situ sap flow observations from a high latitude temperate deciduous forest 172 ecosystem at Soroe in Denmark, to assess the impacts of f_{diff} on GPP and ET. At this site, 73.54% of all days are 173 non-clear ($f_{\text{diff}} > 50\%$). This percentage is higher than the global average level (ca. 50%, Kanniah et al., 2012). 174 The specific objectives are: (1) to evaluate how f_{diff} independently and jointly with other biophysical constraints affects daily carbon and water fluxes in a deciduous forest; (2) to incorporate f_{diff} as a biophysical constraint into 175 remote sensing based 'top-down' models to improve GPP and ET simulations. This study provides insights on 176 the relative contribution of f_{diff} to the total variability on daily carbon and water fluxes encountered over multiple 177 178 years using both statistical path analysis and global sensitivity analysis of 'top-down' models. It also explores 179 potential mechanisms increasing LUE, WUE and evaporative fraction (EF) under diffuse conditions for high 180 latitude ecosystems.

181

182 2. Study site and data

183 A Danish temperate deciduous beech forest site (Soroe on Zealand, Denmark, 55°29'N, 11°38'E) has been 184 selected to evaluate the impacts of f_{diff} on the ecosystem carbon and water fluxes. The Soroe flux site has longterm records of eddy covariance fluxes since 1996, diffuse / total PAR measurements during the period from 2004 to 2013, and sap flow data from 2009 to 2011. The mean annual precipitation is 564 mm and the mean annual temperature is 8.5 °C. The dominant tree species is European beech (*Fagus sylvatica* L.) with approximately 20% conifers, mainly Norway spruce (*Picea abies* (L.) Karst.) and European larch (*Larix decidua* (Mill.)) (Wu et al., 2013). Leaf area index (LAI) peaks at 4-5 m²·m⁻². Soil was classified as Alfisols or Mollisols with 10-40 cm deep organic layers. Details of this site are reported in Pilegaard et al. (2001) and Pilegaard et al. (2011).

Eddy covariance and micrometeorological observations and satellite data from the Moderate Resolution Imaging 192 Spectroradiometer (MODIS) onboard of TERRA were used. The diffuse and total incoming PAR were measured 193 194 by the Delta-T BF3 sensor. Eddy covariance and micrometeorological observations include GPP, ET, Rn-G, incoming longwave radiation (LW_{in}), outgoing longwave radiation (LW_{out}) and incoming shortwave radiation 195 (SW_{in}), air temperature (Ta), vapor pressure deficit (VPD) and soil water content (SWC). More details on this 196 197 dataset can be found in Wu et al. (2012). The initial half-hourly observations were downloaded from the Fluxnet 198 database (https://fluxnet.ornl.gov/), filtered by quality control flags and energy closure errors, and aggregated into daily values. Flag quality controlled GPP and ET observations spanning the period from 2002 to 2012 were 199 200 used for analysis and modeling. For ET, observations with negative sensible heat flux, latent heat flux and net radiation were further excluded. Sap flow data were measured continuously for six beech trees during the period 201 202 of 2009-2011 using the stem-heat balance technique (Granier et al., 1985). Averaged data from these six trees 203 were used to represent the ecosystem-scale transpiration and to evaluate the simulated transpiration. Due to 204 technical issues, there were gaps in the sap flow data. The daily ecosystem scale transpiration was only 205 calculated, if more than three tree observations per day are available. Normalized difference vegetation index 206 (NDVI) from the MODIS satellite vegetation index product (MOD13Q1, 16 day composite at 250m resolution 207 L3 product, https://reverb.echo.nasa.gov/) was downloaded to infer the vegetation phenology and to retrieve LAI 208 dynamics from 2002 to 2012. The initial 16-day synthetic data were further smoothed by the Savitzky–Golay filter in order to reduce the impacts of clouds and then interpolated into daily data by the spline algorithm (Chen 209 et al., 2004). Further, LAI was obtained from NDVI by the locally empirical relationship 210 LAI = 0.001306e^{9.241NDVI} from Boegh et al. (2009). Both LAI from MODIS and in-situ measurements by LAI-211 2200C Plant Canopy Analyzer (LI-COR Inc., Lincoln, NE, USA) were shown in Figure 1. In general, LAI from 212 MODIS NDVI captured the seasonal dynamics of vegetation. The peak values are 4-5.5 m²·m⁻² and these match 213 214 the in-situ measurements and previous studies (Pilegaard et al. 2011; Wu et al., 2013).



Figure 1. The seasonal variation of LAI derived from MODIS NDVI in Soroe from 2002 to 2012 (continuous line). Dots are in-situ LAI measurements from LAI-2200C and the error bar shows the standard deviation.

Surface temperature, Ts, was calculated from in-situ incoming and outgoing longwave radiation based on the
Stefan–Boltzmann law, as in Eq. (1). The surface emissivity was estimated from NDVI (Van de Griend and Owe,
1993, Eq. 2).

$$\varepsilon \cdot \sigma \cdot Ts^4 = LW_{out} - (1 - \varepsilon) \cdot LW_{in}$$
 (Eq. 1)

222
$$\epsilon = \begin{cases} 0.986 & (NDVI > 0.608) \\ 1.0094 + 0.047 \cdot \ln(NDVI) & (0.131 < NDVI < 0.608) \\ 0.914 & (NDVI < 0.131) \end{cases}$$
(Eq. 2)

223 Where LW_{out} is longwave outgoing, LW_{in} is longwave incoming. ε is the surface emissivity and σ is the Stefan-224 Boltzmann constant (5.670367×10⁻⁸ kg·s⁻³·K⁻⁴). NDVI is from MODIS products. 0.986 is emissivity for dense 225 vegetation (NDVI > 0.608) and 0.914 is emissivity for bare soil (NDVI < 0.131).

 f_{diff} is the ratio between the observed diffuse PAR and the total PAR at the ground (Table 1). It is highly 226 227 correlated with the atmospheric transmission. Several indices could be used to infer f_{diff} (Butt et al., 2010). 228 Among them, the cloudiness index (CI), which is one minus the ratio between the observed PAR at ground and 229 PAR at the top of atmosphere (TOA) (Table 1), was often used to represent f_{diff} (Orgill and Hollands, 1977; Butt et al., 2010; Wang et al., 2015). The advantage of CI is that it only requires one measurement, the total PAR at 230 the surface. The PAR at TOA could be calculated based on the time and the location on the Earth. In order to 231 232 determine f_{diff} , ground measurements of both diffuse and total PAR are required. The CI approach was more favorable to be used in places without total and diffuse PAR measurements. In order to make this study more 233 applicable for other regions, CI was adopted as a proxy of f_{diff} to assess its impacts on carbon and water exchange. 234 Additionally in our data set, f_{diff} is available from 2004, while CI has a longer time series since 2002. To identify 235 236 the difference between CI and f_{diff} , these two indices were compared through statistical correlation and modeling 237 tests.

238

239 3. Methods

240 First, statistical analysis was conducted for in-situ eddy covariance and micrometeorological variables from 2002 to 2012 to identify the relationship between CI (a proxy for f_{diff}) and observed daily GPP, ET, LUE, EF and 241 WUE. Then, a joint GPP and ET 'top-down' parsimonious model was used to simulate daily GPP and ET. The 242 model is based on the remote sensing LUE GPP model (Potter et al., 1993; Monteith et al., 1972) and Priestley-243 244 Taylor Jet Propulsion Laboratory ET model (PT-JPL, Fisher et al., 2008). In this model, the same biophysical 245 constraints were used to reduce GPP and ET from potential to actual values. Model accuracy was compared for the cases with and without considering CI. A global sensitivity analysis (GSA, Saltelli et al., 2010) was used to 246 quantify the sensitivity of GPP and ET to f_{diff} . Both statistical analysis and model based GSA provide estimates 247 of the independent and joint effects of CI on environmental variables. Their results were compared in order to 248 249 thoroughly understand effects of CI and related meteorological variables on the carbon and water exchange. 250 Finally, to check the difference between CI and f_{diff} , these two indices were compared through statistical analysis 251 and modeling tests.

252 3.1 Statistical analysis

To qualitatively explore the responses of GPP and ET to different levels of CI, relationships for GPP vs. PAR, 253 254 ET vs. Rn-G, ET vs. PAR were analyzed under predominantly diffuse (CI>0.66) or direct (CI<0.33) radiation conditions. We chose 0.66 and 0.33 as thresholds to have equal intervals between 0 and 1, following the 255 thresholds adopted in Davin and Seneviratne (2012). It has been shown in other ecosystems that CI can covariate 256 257 with vegetation phenology obscuring the actual contribution of CI to GPP (Williams et al., 2016). To address 258 that and control for the phenology effect on GPP and ET, we compared LUE, EF and WUE under diffuse or 259 direct radiation conditions for different levels of NDVI, to make sure that they reflect the same phonological state. Afterwards, path analysis was used to quantitatively assess these relationships. These results informed 260 parsimonious model design by identifying the most important drivers of GPP and ET in this ecosystem. The 261 statistical tests were performed in a significance level of p < 0.05 (1.96×Standard Error). 262

263 Path analysis is a multiple regression technique that considers the covariance among variables. It is mainly used for variables that are highly correlated (Li, 1975) e.g. PAR, Ta and VPD. This method has been applied to 264 evaluate environmental controls on carbon exchange in various ecosystems (Bassow & Bazzaz, 1998; Huxman 265 et al., 2003; Wu et al., 2016). It assumes that the correlation between variable i and dependent variable y can be 266 decomposed into direct and indirect effects. Where the direct effect means that input variable i directly affects 267 268 output variable y. The direct value is also called path value and is the standardized partial regression coefficient. 269 The indirect effects consider how variable i influences another variable j ($j \neq i$) which in turn affects the output y. 270 For example, a direct effect of CI on GPP will reflect the net change in GPP due to solely CI, while the rest of 271 variables (e.g. PAR or air temperature) are fixed. However, the indirect effect will reflect how CI influences 272 other environmental variables (e.g. PAR or air temperature) and in turn, these variables influence GPP. These 273 direct and indirect effects represent the relative strength of a given relationship. Eq. 3 shows the formulas for this 274 decomposition.

$$r_{i,y} = r_{i,1}P_{1,y} + r_{i,2}P_{2,y} + \dots + r_{i,i}P_{i,y} + \dots + r_{i,n}P_{n,y}$$
(i=1, 2, 3, ..., n) (Eq. 3)

Where $r_{i,y}$ is the correlation coefficient between input variable i and ouput y. It decomposes into the direct effect $P_{i,y}$ and indirect effects $r_{i,n}P_{n,y}$ (n \neq i). $P_{i,y}$ is the direct effect from input variable i to output variable y. $r_{i,n}$ is the correlation coefficient between the variable i and variable n. $r_{i,n} * P_{n,y}$ (n \neq i) are the indirect effects. The indirect effect quantifies the effect of one variable on another variable, which in turn affects the dependent variable.

To provide a quantitative assessment of the contribution of diffuse/direct PAR to the daily variability of carbon and water fluxes over 11 years and its interactions with other environmental variables, path analysis was performed considering the effect of various environmental factors on the target variables at the daily time scale from 2002 to 2012. Besides CI, environmental factors include Ta, PAR, Rn-G, LAI, VPD and SWC. Target variables are GPP, λET, LUE, EF and WUE.

287 3.2 Joint Gross Primary Productivity and Evapotranspiration model

To simulate the effects of f_{diff} on GPP and ET, a joint LUE GPP and PT-JPL ET model was used. Both 288 289 approaches estimate GPP or ET under potential conditions and then the potential values are down-regulated by 290 the same biophysical constraints reflecting multiple limitations or stresses. These constraints can be derived from 291 remote sensing and atmospheric data (McCallum et al., 2009; Garcia et al, 2013). The LUE GPP model is 292 recognized as a robust method to estimate GPP across various ecosystems and climate regimes (McCallum et al., 293 2009). The PT-JPL ET model has been demonstrated as one of best-performing global remote sensing ET 294 algorithms in multi-algorithm inter-comparisons (Chen et al., 2014; Ershadi et al., 2014; Vinukollu, Meynadier 295 et al., 2011; Vinukollu, Wood et al., 2011; Michel et al., 2016; Miralles et al., 2016). Therefore, these two 'topdown' GPP and ET models were selected for this study. 296

297 Most widely used LUE models e.g. CASA (Potter et al., 1993) or the MODIS algorithm (Running et al., 2004) 298 are based on the assumption that plants optimize canopy LUE or whole canopy carbon gain per total PAR absorbed (Monteith et al., 1972). They have common features to estimate GPP: (1) ecosystem GPP is directly 299 300 related to absorbed PAR (APAR) through LUE, and (2) LUE may be reduced below its theoretical potential value by environmental stresses such as low temperature or water shortage (Landsberg, 1986). The general form 301 of the LUE GPP model used in this study is shown in Eq. (4) and it is partly based on the Carnegie-Ames-302 303 Stanford-Approach (Potter et al., 1993) with improvements by including constraints to account for fraction of the 304 canopy that is photosynthetically active vegetation (Fisher et al., 2008).

305
$$GPP = \varepsilon_{max} \cdot PARc \cdot f_a \cdot f_M \cdot f_{Ta} \cdot f_{VPD} \cdot f_{SWC} \qquad Eq. (4)$$

Where GPP is the gross primary productivity (g·C·m⁻²·d⁻¹). ε_{max} is the maximum LUE (g·C··MJ⁻¹). PARc is the 306 daily photosynthetically active radiation (PAR) (MJ·m⁻²·d⁻¹) intercepted by the canopy and it is calculated based 307 on the extinction of PAR within the canopy using the Beer Lambert law (Table 1). f_q is the green canopy 308 309 fraction indicating the proportion of active canopy. f_M is the plant moisture constraint. f_{Ta} is the air temperature 310 constraint reflecting the temperature limitation of photosynthesis. f_{VPD} is the VPD constraint reflecting the 311 stomatal response to the atmospheric water saturation deficit. f_{SWC} is the soil moisture constraint on 312 photosynthesis. All these constraints range from 0 and 1 and represent the reduction of maximum GPP under 313 limiting environmental conditions. For more details, see Table 1 and Fisher et al., (2008).

The Priestley-Taylor Jet Propulsion Laboratory model (PT-JPL, Fisher et al., 2008) is based on the Priestley and Taylor (1972) equation for potential evapotranspiration, and incorporates eco-physiological variables to downregulate potential evapotranspiration to actual evapotranspiration. PT-JPL is a three source evapotranspiration model, which includes wet surface evaporation (Ei), transpiration (Ec) and soil evaporation (Es), as descripted in equations (5-8).

319 $\lambda ET = \lambda Ei + \lambda Ec + \lambda Es$ Eq. (5)

320
$$\lambda \text{Ei} = f_{wet} \cdot \alpha \Delta / (\Delta + \gamma) \cdot \text{Rnc}$$
 Eq. (6)

321
$$\lambda \text{Ec} = (1 - f_{wet}) \cdot f_g \cdot f_M \cdot f_{Ta} \cdot \alpha_c \Delta / (\Delta + \gamma) \cdot \text{Rnc} \qquad \text{Eq. (7)}$$

322
$$\lambda Es = f_{SWC} \cdot \alpha \Delta / (\Delta + \gamma) \cdot (Rns - G)$$
 Eq. (8)

Where λET is the latent heat flux for total evapotranspiration (W·m⁻²), λEi is evaporation of intercepted water 323 $(W \cdot m^{-2})$, λEc is transpiration $(W \cdot m^{-2})$, and λEs is evaporation of soil water $(W \cdot m^{-2})$. The quantity f_{wet} is the 324 relative surface wetness to partition the evapotranspiration from the intercepted water and canopy transpiration 325 (Fisher et al., 2008). The symbols f_q , f_M , f_{Ta} and f_{SWC} denote biophysical constraints and have the same 326 327 meaning as in Eq. 4. f_{wet} They vary from 0 to 1 to account for the relative reduction of potential λ ET under 328 limiting environmental conditions. Rnc and Rns are the net radiation for canopy and soil, respectively. The 329 partitioning of PAR and net radiation between canopy and soil is calculated following the Beer-Lambert law 330 (Table 1). G is the ground heat flux. Δ is the slope of saturation-to-vapor pressure curve. γ is the psychrometric constant. α is an empirical ratio of potential evapotranspiration to equilibrium potential evapotranspiration (PT 331 332 coefficient) replacing the atmospheric demand and surface resistance effects in the Penman-Monteith ET equation. Here for λEi and λEs , α is equal to 1.26. This is also the suggested value for the PT-JPL model (Fisher 333 et al., 2008). α_c is the coefficient for λEc and it is the only parameter in the model that requires calibration. 334

In order to make the models parsimonious and robust, only those constraints/variables having significant relationships with GPP and λ ET were included in the LUE and PT-JPL models. Table 1 shows the detailed information on the model constraints and parameters for LUE and PT-JPL models.

Table 1. Model parameters and equations. SZA is the sun zenith angle. The extinction coefficients for PAR (k_{PAR}) and for net radiation (k_{Rn}) were equal to 0.5 and 0.6, respectively (Ross, 1976; Impens & Lemur, 1969; Fisher et al., 2008), RH is the relative humidity.

Parameter	Description	Equation	Reference
f_g	Green canopy fraction	$f_g = f_{APAR} / f_{IPAR}$	Fisher et al., 2008
f _м	Plant moisture constraint	$f_M = f_{APAR} / \max(f_{APAR})$	Fisher et al., 2008
f_{Ta}	Plant temperature constraint	$f_{Ta} = 1.1814 \cdot [1 + e^{0.3(-T_0 - 10 + Ta)}]^{-1} [1 + e^{0.2(T_0 - 10 - Ta)}]^{-1}$	Potter et al., 1993
<i>f_{swc}</i>	Soil moisture constraint	$f_{SWC} = \frac{SWC - SWC_{\min}}{SWC_{\max} - SWC_{\min}}$	Fisher et al., 2008
$f_{\scriptscriptstyle VPD}$	Vapor pressure deficit constraint	$f_{VPD} = 1/(1 + VPD/D0)$	Lohammar et al., 1980
f _{wet}	Relative surface wetness	$f_{wet} = \mathrm{RH}^4$	Fisher et al., 2008
f _{ci}	Cloudiness index constraint	$f_{ci} = 1 - rac{\mathrm{CI} - \mathrm{CI}_{\min}}{\mathrm{CI}_{\max} - \mathrm{CI}_{\min}}$	This study
f _{di}	Fraction of diffuse PAR constraint	$f_{di} = \frac{f_{diff} - min(f_{diff})}{max(f_{diff}) - min(f_{diff})}$	This study
PARc	PAR intercepted by the canopy	PARc = PAR - PARs	Ruimy et al., 1999
PARs	PAR for the soil	$PARs = PAR \cdot e^{\frac{-k_{PAR} \cdot LAI}{\cos(SZA)}}$	Ruimy et al., 1999
Rnc	Net radiation for the canopy	Rnc = Rn - Rns	Fisher et al., 2008
Rns	Net radiation for the soil	$Rns = Rn \cdot e^{\frac{-k_{Rn} \cdot LAI}{\cos(SZA)}}$	Fisher et al., 2008

LAI	Leaf area index	$LAI = 0.001306e^{9.241NDVI}$	Boegh et al., 2009
<i>f_{APAR}</i>	Fraction of PAR absorbed by green vegetation cover (SAVI: soil adjusted vegetation index)	SAVI = 0.45 NDVI + 0.132 $f_{APAR} = 1.4 SAVI - 0.05$	Fisher et al., 2008
<i>f</i> _{IPAR}	Fraction of PAR intercepted by total vegetation cover	$f_{\rm IPAR} = 1.0 \text{ NDVI} - 0.05$	Fisher et al., 2008
CI	Cloudiness index	$CI = 1 - PAR_{obs} / PAR_{TOA}$	Spitters et al., 1986
f _{diff}	Fraction of diffuse PAR	$f_{diff} = PAR_{diffuse} / PAR_{total}$	Spitters et al., 1986
То	Optimum plant growth temperature	Ta at max{PAR $\cdot f_{APAR} \cdot \text{Ta/VPD}$ 16.51 °C for this study	Fisher et al., 2008
D0	Empirical coefficient for VPD	15 hPa	Leuning et al., 1995



342 3.3 Incorporating diffuse fraction into the joint GPP and ET model

Previous studies have improved the LUE GPP models by considering the impacts of CI on LUE (e.g. Turner et al., 2006; Wang et al., 2015; Wu et al., 2016). For this study, we used a similar approach to modify Eq. (4) to incorporate the CI constraint into the GPP model, as Eq. (9). Moreover, ε_{max} now represents the maximum LUE under totally diffuse radiation conditions instead of the maximum value for all sky conditions.

In the PT-JPL model, the PT coefficient (α) represents the atmospheric demand and the surface resistance for ET. Therefore, similar to the maximum LUE ε_{max} in the GPP model, this study incorporated CI into ET via changes in the PT coefficient (α_c), which reflects the opening of stomata and stomatal conductance.

$$GPP = (1 - \mu \cdot f_{ci}) \cdot \varepsilon_{\max} \cdot PARc \cdot f_g \cdot f_M \cdot f_{Ta} \cdot f_{VPD} \cdot f_{SWC}$$
 Eq. (9)

$$\lambda \text{Ec} = (1 - \tau \cdot f_{ci}) \cdot (1 - f_{wet}) \cdot f_g \cdot f_{Ta} \cdot f_M \cdot \alpha_c \Delta / (\Delta + \gamma) \cdot \text{Rnc} \qquad \text{Eq. (10)}$$

Where f_g , f_M , f_{Ta} , f_{VPD} , f_{SWC} , f_{wet} have the same meaning as Eq. (2) and (7). PARc and Rnc are the PAR and Rn intercepted by the canopy, respectively. $(1 - \mu \cdot f_{ci})$ and $(1 - \tau \cdot f_{ci})$ are the subtractive formulas to represent the fraction of diffuse PAR constraints for GPP and ET, respectively. μ indicates an overall sensitivity of GPP to CI. τ reflects the sensitivity of λ Ec to CI. In these approaches, f_{ci} can also be replaced by f_{di} , in order to drive the model with f_{diff} instead.

- 357
- 358 3.4 Model calibration and validation

The LUE GPP model version without considering diffuse light (Eq. 2) had only one parameter, ε_{max} (maximum LUE) to be optimized or adjusted to the vegetation type. In the new diffuse/direct model version (Eq. 9), an additional parameter μ (the sensitivity of GPP to CI) needs to be optimized as well. In the initial PT-JPL ET model (Eq. 7), the α_c parameter was optimized, while in the modified PT-JPL ET model (Eq. 10), τ (the sensitivity of λ Ec to CI) needs to be optimized additionally. Table 2 shows the details on these calibrated parameters.

	Models	Parameter	Without CI (Eq. 4 and 7)	With CI (Eq. 9 and 10)	Range	Optimized value (Without CI / with CI)
	GPP	ε_{max}	Maximum LUE (g·C·m ⁻ ² ·MJ ⁻¹)	Maximum LUE under total diffuse PAR conditions (g·C·m ⁻² ·MJ ⁻¹)	0~5	2.97 / 4.29
	μ		GPP sensitivity to f_{CI} (dimensionless)	-1~1	0.46	
	ET	α _c	PT coefficient for the canopy (dimensionless)	PT coefficient for the canopy under total diffuse PAR conditions (dimensionless)	1~3	1.32 / 2.60
	τ		λ Ec sensitivity to f_{CI} (dimensionless)	-1~1	0.65	

365 Table 2. The calibrated parameters for the joint GPP and ET model

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The Monte Carlo method was used to optimize these model parameters, with RMSE (root mean square error) as the objective function. The parameter values were sampled 20,000 times with uniform distribution within their corresponding ranges and these 20,000 parameter sets were used to run models. Odd years were used for calibration and even years were for validation. The best-fit parameter set was chosen.

To compare model simulation performances with and without CI information, the Root Mean Square Error 371 372 (RMSE, Eq. 11), Correlation Coefficient (R, Eq. 12), Bias (Eq. 13), unbiased Root Mean Square Error (ubRMSE, Eq. 14) and Standard Deviation (STD, Eq. 15) were used. Taylor diagrams (Taylor, 2001), were used to present 373 these three complementary statistics CC, Normalized STD (NSTD, as Eq. 16) and Normalized ubRMSE 374 (NubRMSE), which have a triangle-cosine-law-like relationship, as Eq. (17). The radial distance stands for the 375 NSTD and the angle in the polar plot represents R. The reference point located on the X-axis with R=1, NSTD=1 376 and NubRMSE=0 is the observation. The distance from the simulation point to the reference point means the 377 378 NubRMSE of simulations and it is the integrated performance for the simulation.

379
$$RMSE = \sqrt{\sum_{i=1}^{N} (sim_i - obs_i)^2 / N}$$
(11)

380
$$R = \frac{\sum_{i=1}^{N} (\sin_i - \overline{\sin})(obs_i - \overline{obs})}{\sqrt{\sum_{i=1}^{N} (\sin_i - \overline{\sin})^2} \times \sqrt{\sum_{i=1}^{N} (obs_i - \overline{obs})^2}}$$
(12)

381
$$BIAS = \sum_{i=1}^{N} (sim_i - obs_i) / N$$
(13)

382
$$ubRMSE = \sqrt{\sum_{i=1}^{N} [(sim_i - \overline{sim}) - (obs_i - \overline{obs})]^2 / N}$$
(14)

$$STD = \sqrt{\sum_{i=1}^{N} (\sin - \overline{sim})^2 / N}$$
(15)

$$NSTD_{sim} = STD_{sim} / STD_{obs}$$
(16)

385
$$NubRMSE_{obs,sim}^{2} = NSTD_{obs}^{2} + NSTD_{sim}^{2} - 2NSTD_{obs}NSTD_{sim} \cos CC_{obs,sim}$$
(17)

Where sim is the simulation, obs is the observation, N is the total number, \overline{sim} is the average of the simulation, and \overline{obs} is the average of the observation.

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389 3.5 Global sensitivity analysis

390 The Sobol' method (Sobol' et al., 1990) is one of the most commonly used global sensitivity analysis (GSA) methods. It is based on ANOVA (analysis of variance) decomposition and it allows calculating the sensitivity of 391 392 coupled input forcing. The Sobol' method provides not only the first order sensitivity for each forcing factor but 393 can also quantify interactions among forcing factors. The first order sensitivity quantifies the independent 394 contribution from each forcing to the output variable, while the second order quantifies the interactions between 395 each two forcing factors to the output variable. In our study, we aimed to identify the sensitivity of GPP and λET 396 to f_{ci} , in relation with other major environmental variables and assess if the model approach could pick the same 397 sensitivities embedded in the dataset that will be captured by the path analysis. The variances of the terms in the ANOVA decomposition are estimated the following equations (Saltelli et al., 2010): 398

$$V(Y) = \sum_{i=1}^{n} V_i + \sum_{i\leq j}^{n} V_{ij} + \dots + \sum_{1\leq \dots \leq n}^{n} V_{1\dots n}$$
(18)

400 Where V_i represents the first order effect for each factor X_i ; V_{ij} stands for the second order effect for X_i , X_j ; and 401 $V_{1...n}$ is the nth order effect for $X_1, ..., X_n$.

402 The first order sensitivity index S_i can be calculated by

403
$$S_i = V_i / V(Y) = V[E(Y | X_i)] / (V(Y))$$
 (19)

404 And the second-order sensitivity index S_{ij} can be calculated by

405
$$S_{ij} = V_{ij}/V(Y) = (V[E(Y | X_i, X_j)] - V_i - V_j)/(V(Y))$$
(20)

406 In general, the total sensitivity index can be defined as:

$$S_i^{\text{tot}} = E(V(Y|X_{\sim i}))/V(Y)$$
(21)

408 Where S stands for different order sensitivity index, V means the variance for different variables, E is the 409 expectation, and ~i refers to all of the inputs except input i.

410 The kernel density sampling method was applied to sample the input data set for sensitivity analysis. The 411 advantage to use the kernel density sampling method is that it could resemble the distribution of sampled data set. 412 According to the kernel density distribution of each model input, 20,000 samples will be generated to assess the 413 model input sensitivity.

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- 415 4. Results and discussion
- 416 4.1 Statistical analysis

First, we explored responses of daily GPP and λ ET to diffuse/direct radiation conditions while controlling for 417 radiation levels. Daily GPP and λET from the eddy covariance flux tower were compared between 418 419 predominately diffuse PAR conditions (CI>0.66) and predominately direct PAR conditions (CI<0.33) over the 420 period from 2002 to 2012. Direct comparison for diffuse and direct radiation conditions could involve data from different days of the year with different phenology conditions and this comparison could exaggerate the actual 421 422 diffuse fertilization effects (Williams et al., 2016). In order to compare the responses of daily GPP and λ ET with the phenology background, comparison in Figure 2 (a-c) was conducted with NDVI more than 0.75. Figure 2 (a-423 c) shows the evolution of GPP and λ ET as a function of PAR and Rn-G with predominantly diffuse and direct 424 PAR conditions, respectively. There is a clear distinction of the response of GPP to PAR between diffuse and 425 direct PAR conditions. As shown in Figure 2 (a), with PAR increasing, GPP increased significantly in the diffuse 426 427 PAR conditions, while GPP increased slowly in the direct PAR conditions. The slope of the response curve 428 represents LUE. This suggests that LUE in the diffuse PAR is higher than that in the direct radiation. The 429 evolution of λ ET as a function of PAR or Rn-G in diffuse and direct radiation conditions is shown in the Figure 430 2 (b) and (c), respectively. Similar to the response of GPP, λET is higher under predominately diffuse PAR conditions for the same level of PAR or Rn-G. However, the increase in λET is less obvious than that in GPP 431 and this leads to the increase of WUE. These results are in agreement with a study on a deciduous temperate 432 433 forest ecosystem of central Germany, which found that the diffuse/direct radiation could increase the ecosystem WUE (Knohl and Baldocchi, 2008). The slope of the response curve in Figure 2 (c) represents the evaporative 434 fraction (EF), the ratio of between λ ET and Rn-G. This suggests that under diffuse PAR, higher photosynthesis 435 436 rates lead to higher λET and higher EF. It should be also noticed that the difference for response curves in λET 437 vs. PAR is clearer than that in λ ET vs. Rn-G. PAR does not include the information on the longwave radiation, while Rn-G contains the longwave radiation components. That indicates that longwave radiation components 438 439 induced the difference between λET vs. PAR and λET vs. Rn-G.

440 To further explore the responses of daily LUE, EF and WUF to diffuse/direct radiation conditions, LUE, EF and 441 WUE were compared with various levels of NDVI under predominately diffuse and direct PAR conditions, as 442 shown in Figure 2 (d-f). In general, there is a significant difference for the response curves of diffuse and direct PAR conditions in LUE, EF and WUE. This indicates the ecosystem responds differently to the diffuse and 443 444 direct PAR conditions. With higher levels of NDVI, the difference of LUE, EF and WUE between predominately diffuse conditions and predominately direct radiation conditions becomes more significant. This 445 446 indicates diffuse PAR has stronger effects in high NDVI conditions. This is in agreement with the findings that 447 in the ecosystem with high LAI, the diffuse fertilization effects are stronger (Alton et al., 2007).



Figure 2. The response of daily GPP, λ ET, LUE, EF and WUE to diffuse and direct radiation conditions during 450 the period from 2002 to 2012. The thresholds for predominantly diffuse and direct conditions are defined as CI 451 452 above 0.66 and below 0.33, respectively. (a) GPP as a function of PAR for predominantly diffuse light condition (CI>0.66) and predominantly direct light conditions (CI<0.33). (b) λ ET as a function of PAR for predominantly 453 454 diffuse and direct light conditions. (c) λ ET as a function of Rn-G for predominantly diffuse and direct conditions. 455 (d) LUE as a function of NDVI for predominantly diffuse and direct conditions. (e) EF as a function of NDVI for predominantly diffuse and direct conditions. (f) WUE as a function of NDVI for predominantly diffuse and 456 direct conditions. The points represent the mean value for specific PAR interval and the error bar represent the 457 458 significance level at p < 0.05 (1.96×Standard Error). To exclude the influence from phenology, comparison in (a-c) were conducted with NDVI more than 0.75. 459

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The results of the path analysis are shown in Table 3-5. Table 3 presents the total effects (correlation coefficients) 461 462 among these important variables. CI has positive correlation with LUE, EF, EF* and WUE, while it negatively 463 correlates with GPP, λ ET and λ ET*. λ ET* and EF* are the observations with LAI greater than 2 and VPD larger than 3.5 hPa. LAI greater than 2 corresponds to the growing season, while VPD larger than 3.5 hPa is associated 464 with lack of precipitation and limited evaporation of intercepted water. Both Ta and LAI positively correlate 465 with GPP, LUE, λET, λET*, EF, EF* and WUE. PAR also has positive correlation with GPP, LUE and WUE. 466 Rn-G has positive correlations with λET , λET^* and EF. This indicates that the ecosystem dynamics are 467 468 controlled by temperature and radiation. During the growing season, there is a negative correlation between VPD and EF*, reflecting stomatal control of transpiration. This can be seen also in the negative correlation of VPD 469 470 with EF*. Even though there are parts of the year when the ecosystem is water controlled, the overall dynamics 471 are controlled by energy and temperature. This is supported by the negative correlation of SWC with GPP, LUE, 472 λ ET and EF. The relation is only positive during high VPD and growing season periods (λ ET*). In water-limited

473 ecosystems or situations, EF usually has negative correlation with VPD while SWC positively correlates with
474 GPP, LUE, λET and EF.

			/		υ			
Total effects	CI	Та	PAR	Rn-G	LAI	VPD	SWC	
 GPP	-0.47	0.78	0.84	/	0.84	0.72	-0.31	
LUE	0.11	0.71	0.28	/	0.73	0.27	-0.36	
λΕΤ	-0.29	0.73	/	0.76	0.77	0.64	-0.26	
λET^*	-0.21	0.15	/	0.68	0.34	0.29	0.25	
EF	0.12	0.59	/	0.11	0.63	0.21	-0.31	
EF*	0.44	0.12	/	-0.36	0.20	-0.19	0.09	
WUE	0.08	0.25	0.04	-0.05	0.22	0.03	-0.20	

475 Table 3. The total effects (correlation coefficients) from environmental factors to target variables

476 '/' means not the input for the correlation test.

478 The total effects were further decomposed into direct and indirect effects using path analysis. Table 4 shows the direct effects of environmental variables on the target variables. GPP, λET and WUE are mainly controlled by 479 radiation (either PAR or Rn-G). Once normalizing for different radiation as in LUE or EF variables the most 480 important factor was Ta. CI had significant positive direct effects on all target variables GPP, LUE, λ ET, EF and 481 482 WUE. That means that an increase in CI while maintaining the rest of considered variables fixed will produce a 483 net increase in GPP, LUE, ET and EF. Based on this when incorporating CI into the models, we should consider that CI will increase GPP and ET (see Eq. (9) and (10)). Considering the whole year, the effects of CI on GPP, 484 LUE and WUE are stronger than on λ ET and EF. It is possible that the effect of CI on transpiration is masked by 485 486 evaporation from soil and intercepted water, which are insensitive to CI. When considering λET^* and EF^* , to 487 minimize the effect of evaporation of intercepted water and soil water, CI had stronger direct effects. This agrees with the land surface modeling results by Davin and Seneviratne (2012), which show that CI mainly influences 488 transpiration and has limited impacts on evaporation from the intercepted water and soil. This finding further 489 supports our modeling approach, which incorporates CI into the transpiration module only. Variables related to 490 491 light harvesting by canopies (PAR and LAI) were the dominant factors regulating GPP. The top soil moisture (SWC) has very limited effects on GPP. For LUE, Ta was the important factor with a positive correlation, which 492 emphasizes the sensitivity of this ecosystem to temperature. After Ta, LAI had also a positive effect while VPD 493 reduced LUE. CI ranked as the fourth most important factor to influence LUE with a positive response. Similar 494 to GPP, SWC had very weak effects on LUE. It is possible that the deep rooting system of the beech forest 495 enables sufficient water supply even though water in the top soil is depleted (Wu et al., 2012). From these 496 497 findings, we can conclude that the GPP and LUE of this high latitude ecosystem are controlled by radiation and temperature. For λ ET, CI had a weaker but still significant influence considering the whole year. As expected, 498 499 CI had stronger impacts on λET^* during the growing season and in periods of high VPD. When normalizing λET 500 by the available energy (Rn-G), Ta was the dominant factor with a positive effect on EF, followed by LAI and SWC. During the growing season, CI became the major controlling factor for EF*. 501

In our site, after accounting for energy, Ta and CI are the main factors to influence WUE. WUE increased in response to increases in Ta and CI. This is different from water-limited ecosystems, where increases in Ta tend to decrease WUE (Stroosnijder et al., 2012) and diffuse PAR tends to compensate for this effect (Gu et al., 2002;

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805 Rocha et al., 2004). However, higher VPD reduces GPP, LUE and λ ET and EF due to the strong stomatal control.

506 These findings are helpful to refine parsimonious models for the simulation of GPP and ET. For instance, SWC

507 of the top soil could not be the necessary input for GPP simulation, since it has weak impacts on GPP.

-					5				
-	Direct effects	CI	Та	PAR	Rn-G	LAI	VPD	SWC	
-	GPP	0.19	0.07	0.82	/	0.44	-0.17	0.02	
	LUE	0.31	0.58	0.20	/	0.47	-0.42	-0.02	
	λΕΤ	0.03	0.14	/	0.51	0.50	-0.05	0.11	
	λET^*	0.12	0.01	/	0.67	0.25	0.00	0.18	
	EF	0.02	0.70	/	-0.35	0.46	-0.30	0.08	
	EF*	0.25	0.23	/	-0.29	0.20	-0.14	0.22	
	WUE	0.32	0.30	0.76	-0.46	-0.12	-0.24	-0.20	

508 Table 4. The direct effects from environmental factors to target variables

509 ET* and EF* are for conditions (LAI > 2 and VPD > 3.5 hPa). '/' means not the input for path analysis.

511 The indirect effects describe how CI influences ecosystem carbon and water fluxes through other intermediate 512 environmental variables over the whole year or the growing season, as shown in Table 5. Over the year, CI 513 mainly interacts with variables related to the radiation transfer (PAR, Rn-G and LAI) to reduce GPP and λ ET, 514 respectively. After these radiation variables, Ta and VPD have been shown to deliver major indirect effects from 515 CI to GPP, LUE, λ ET and EF. Higher CI over the year decreases Ta, decreasing in turn LUE and EF, but CI also 516 reduces VPD, which has positive effects on LUE and EF. For WUE, the effects of CI mainly go through PAR, 517 Rn-G and VPD. It should be also noticed that the paths through SWC are very weak or not significant.

518 Table 5. The indirect effects from CI through other environmental factors to target variables

Indirect effects from CI via:	Та	PAR	Rn-G	LAI	VPD	SWC
GPP	-0.02	-0.63	/	-0.12	0.11	0.00
LUE	-0.18	-0.16	/	-0.13	0.27	0.00
λΕΤ	-0.03	/	-0.28	-	0.03	-
λET^*	0.00	/	-0.36	0.04	0.00	-
EF	-0.19	/	0.19	-0.05	0.15	-
EF*	-0.03	/	0.14	0.02	0.06	-
WUE	-0.06	0.28	-0.59	-	0.14	-0.01

519 ET* and EF* are for conditions (LAI > 2 and VPD > 3.5 hPa). '/' means not the input for path analysis; '-' 520 means the path analysis is not significant.

522 4.2 Joint Gross Primary Productivity and Evapotranspiration modeling

523 Based on path analysis, a parsimonious GPP and ET model was developed for the site. Since SWC was not a 524 significant factor influencing GPP in the path analysis, the soil moisture constraint was excluded in the GPP and

transpiration modeling but not for soil evaporation. The optimized parameter values increased, as shown in the

Table 2. ε_{max} increased from 2.97 to 4.29 and α_c increased from 1.32 to 2.60. The increase of parameter values

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⁵²¹

is similar to that in Wang et al. (2015). In their study, CI was incorporated into the MODIS LUE algorithm and after optimization, ε_{max} changed from 1.12 to 3.87 in a similar mixed deciduous broadleaf and evergreen needle forest as this study. The value of 2.60 for α_c is also reasonable, considering for forests under totally diffuse radiation conditions. Depending on the land conditions, α_c can reach a value of up to 3.62 in grass and forest (Lhomme, 1997).

532 The model performance of GPP and λ ET with and without CI was compared in Figure 3 during the calibration and validation periods. It can be seen that with CI, the skills of both GPP and ET models improved. During the 533 calibration period, as shown in Figure 3 (a), RMSE of the simulated GPP decreased from 1.55 to 1.34 g·C·m⁻²·d⁻¹ 534 (RMSE reduced 13.25%). R² increased from 0.93 to 0.95. Additionally, the bias was reduced from -7.09% to -535 536 4.14%. During the validation period, as shown in Figure 3 (b), RMSE of simulated GPP reduced from 1.64 to 537 1.45 g·C·m⁻²·d⁻¹ (RMSE reduced 11.68%). R² improved from 0.92 to 0.94. The bias changed from -5.59% to -538 2.58%. Further, we compared the improvement of the simulation performance between the whole period and the growing season. The RMSE of the simulated GPP during the whole period decreased from 1.59 to 1.39 g·C·m⁻ 539 ²·d⁻¹ (RMSE reduced 12.58%) and R² increased from 0.93 to 0.94. While during the growing season, the RMSE 540 dropped from 2.37 to 2.06 g·C·m⁻²·d⁻¹ (by 13.08%) and R² increased from 0.68 to 0.73. It can be seen that there is 541 more improvement in the growing season. 542

543 For λ ET, during the calibration period as shown in Figure 3 (c), when incorporating CI, the RMSE of simulated λ ET from the PT-JPL model decreased from 17.03 to 14.39 W·m⁻² (by 15.50%). R² increased from 0.85 to 0.87. 544 The bias was reduced from 4.94% to 1.08%. During the validation period, shown in Figure 3 (d), the RMSE was 545 reduced from 15.79 to 14.50 W·m⁻² (by 8.16%). \mathbb{R}^2 improved from 0.86 to 0.87, while the bias was reduced from 546 5.20% to 0.09%. We also found for the whole period, the RMSE dropped from 16.09 to 14.44 W·m⁻² (by 547 10.25%), while during the growing seasons, the RMSE decreased from 19.08 to 16.89 W·m⁻² (by 11.47%). Both 548 GPP and λ ET simulation improves when incorporating CI into the models, especially during the growing season. 549 550 The improvement of GPP simulations is more significant than that of λET . This agrees with the results from the 551 statistical path analyses, which showed a higher effect of diffuse PAR on GPP than on λ ET. CI has stronger 552 effects during the growing season (λ ET*) than the effects during the whole period (λ ET). The higher sensitivity of GPP to f_{diff} than for λ ET has been also found in other studies (Mo and Liu, 2001; Knohl and Baldocchi, 2008; 553 554 Oliveira et al., 2011).



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Figure 3. Scatter plots of the simulated and observed GPP (a: calibration, b: validation) and λET (c: calibration,
d: validation). The red dots are simulation with CI and the blue circles are the simulation without CI.

558 To understand under which environmental conditions there is a larger improvement of simulation performance, 559 we stratified the improvement in model errors, $\Delta RSME / MEAN$, by levels of CI, Ta, VPD, LAI, PAR or Rn-G, and also assessed the effect of phenology by considering the day of the year (Figure 4). ARSME is equal to 560 RSME_{without CI} - RMSE_{with CI}. For different levels of CI, the RMSE for both GPP and λ ET simulation decreased 561 after including CI (Figure 4 (a) and (g)) with larger improvements for extreme CI values (e.g. sunny or 562 completely overcast). While for median CI conditions, the improvement is lower. This is because in the 563 simulation without CI, the optimized ε_{max} (Eq. 4) and α_c (Eq. 7) tend to represent median CI conditions, as a 564 compromise that tends to be low in high diffuse fraction conditions and high in low diffuse fraction conditions. 565 In the simulation with CI, ε_{max} (Eq. 9) and α (Eq. 10) were parameterized with CI and the simulation performs 566 well in high and low diffuse radiation conditions. As for temperature shown in Figure 4 (b) and (h), under low 567 568 temperatures the model improvements when incorporating CI are lowest, since both λET and GPP are low. For VPD, both the simulated GPP and λ ET improve at all levels and the improvements of Δ RSME at all levels are 569 similar. However, as for $\Delta RSME / MEAN$ shown in Figure 4 (c) and (i), the λET improvements are larger for 570 high VPD parts and GPP improvements are larger in the low VPD parts. This is due to that with low VPD, 571

572 potential evapotranspiration was partitioned more into evaporation from intercepted water than transpiration, according to Eq. (6) and (7). The model incorporates CI only into the transpiration module. This results in 573 574 limited improvement of λ ET in the low VPD part. However, the GPP improvements of Δ RSME for all levels are 575 similar. The MEAN value of GPP is low. This leads to the high value of $\Delta RSME / MEAN$ in Figure 4 (c). For LAI, as shown in Figure 4 (d) and (j), improvements could be seen with different levels of LAI. For radiation 576 577 (Figure 4 (e) and (k)), the simulation improvements for GPP and λ ET are similar to VPD. This is due to that high VPD and low radiation (low PAR and Rn-G) are concurrent. For different days of the year as shown in Figure 4 578 (f) and (l), the largest improvements occurred in the growing season from May to October. Generally, both GPP 579 and λ ET simulation improvements occur in the growing season, which coincides with higher temperatures and 580 581 larger incoming PAR.



Figure 4. The comparison of modeling performance without and with CI. The y-axis is the Δ RMSE/MEAN. The positive value indicates the simulation improvement, while the negative value means the simulation degradation. (a~f) are for the GPP. (g~l) are for λ ET. (a) and (g) show the modeling improvement with various CI levels. (b) and (h) are for Ta. (c) and (i) are for VPD. (d) and (j) correspond to various LAI levels. (e) and (k) are for different PAR or Rn-G levels. (f) and (l) show the different day of year.

Comparison with sap flow measurements (Figure 5) shows that measurements fit slightly better with the 588 589 simulation with CI (red dots) than the simulation without CI (blue dots). The limited improvement is due to model and observation uncertainties. For example, in the PT-JPL model, the relative surface wetness ($f_{wet} =$ 590 RH⁴) was used to partition the evaporation from the intercepted rainfall and canopy transpiration. This empirical 591 formula might not be transferable without calibration to all sites and f_{wet} may not accurately partition 592 593 evaporation from the intercepted rainfall and canopy transpiration. Additionally, there are uncertainties related to 594 measurements of sap flow and upscaling sap flow data to the ecosystem level. However, the purpose is to compare Ec and sapflow and to check whether there is an improvement of simulated Ec. With the current data 595 set, after incorporating CI into the model, the simulated transpiration improved as shown by the R² increasing 596 from 0.60 to 0.64. This indicates that including CI could improve the simulation of transpiration. By comparing 597 598 the simulated λET of the same dates, the improvement for λET simulation with CI is small. After incorporating 599 CI, the RMSE of λ ET decreased from 13.99 to 13.73 W·m². For R², the simulations with and without CI have the same value of 0.92 (results not shown). This indicates incorporating CI may have limited improvement for 600 601 total ET, but the improvements on Ec are larger.



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Figure 5. The scatterplot to evaluate transpiration (Ec) with sap flow data (blue dots are simulation without CI and red circles are simulation with CI)



In this section, we aim to quantify the influence that CI has on GPP, λ ET and λ Ec and how it varies across 607 different months. The first order sensitivity represents the contribution of model forcing environmental variables 608 609 to GPP, λ ET and λ Ec (Figure 6). For GPP, λ ET and λ Ec, their variations are mainly regulated by the radiation 610 (PAR or Rn). This agrees with the direct effects from path analysis results, as shown in Table 4, indicating that 611 the ecosystem of this flux site is radiation controlled (van Dijk et al., 2005; Lagergren et al., 2008). Moreover, 612 similar to the direct effects from the path analysis, the effects of CI on GPP, λET and λEc are of similar magnitude as those from VPD, air temperature and NDVI (which indicates LAI and phenology). Therefore, the 613 match between the direct effects revealed by path analysis and the first order sensitivity determined with GSA 614 615 confirms that the joint GPP and ET model can capture the major processes in this ecosystem. Furthermore, GSA analysis also shows that sensitivities of GPP, λ ET and λ Ec vary substantially across different months (Fig.5). CI 616 617 had very limited contributions to the variability of GPP, λ ET and λ Ec in winter. The fact that there is a small effect in winter, i.e. when the beech trees are without leaves, is due to the fact that there are ca. 20 % conifers in 618 619 the forest with a small contribution to the annual GPP. However, CI contributed more to the daily variability of GPP, λ ET and λ Ec during the growing season, similar to the findings in our data driven analysis (Table 4) that 620 621 showed that λET^* and EF^* were more sensitive to the CI than λET and EF for the whole years. During the growing season (from May to October), CI contributed to 11.88%, 3.04% and 7.78% of the total variability in 622 623 GPP, λ ET and λ Ec, respectively. The contribution from CI to GPP is the highest, followed by λ Ec. The contribution to λET is less than that to λEc is due to soil evaporation is not sensitive to CI. 624



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Figure 6. The first order Sobol' sensitivity index for the simulated GPP (a), λET (b) and λEc (c) for each month.
The variables are radiation components (PAR for GPP, Rn-G for λET and λEc), CI, VPD, NDVI, Ta and SZA.

The second order sensitivity reveals interactions between the environmental variables and their joint effects on the daily variability of GPP, λ ET and λ Ec. For GPP, as shown in Figure 7 (a), the clearest interaction is between PAR and NDVI, indicating how changes in APAR determine GPP. For λ ET, as shown in Figure 7 (b), the strongest interaction is again for energy and temperature variables: Rn and Ta. Second order sensitivities of λ Ec are shown in Figure 7 (c). Compared to GPP and λ ET, the second order sensitivities of λ Ec are higher indicating more interactions among environmental variables to regulate the transpiration. The highest interaction for λ Ec is between Rn and VPD. CI mainly interacts with radiation (PAR or Rn-G) to control the GPP, λ ET and λ Ec,

reflecting this ecosystem is radiation limited. This result also matches well with the path analysis in Table 5, 635 which CI has highest indirect effects on GPP, λ ET and λ ET* through PAR or Rn-G. NDVI is the second most 636 637 important variable to interact with CI to jointly influence GPP. Similarly, in Table 5, LAI ranked as the second 638 important variable to deliver the indirect effects from CI to GPP. As for λ ET and λ Ec, VPD and NDVI are important variables after Rn-G. Compared to Table 5, the indirect effects from CI through LAI and VPD to λ ET 639 640 and λET^* are also significant. These results indicate that with potentially increasing levels of aerosols and diffuse PAR over the growing season in the future, the WUE, LUE and EF will increase in principle, but the 641 642 magnitude of the enhancement will depend on the interplay between VPD and Ta.



Figure 7. The second order Sobol' sensitivity index for the simulated GPP (a), λET (b) and λEc (c) for the whole year

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647 4.4 Potential mechanism for diffuse PAR to influence GPP and ET of the Sorø beech forest site

648 Several mechanisms have been reported to explain the impacts of f_{diff} on GPP and ET. First, for the same levels of PAR, diffuse PAR penetrates deeper into the canopies than direct PAR, and hence makes the vertical 649 distribution of PAR more even throughout forest canopies. Photosynthesis in the lower part of the canopy will be 650 stimulated by the increased diffuse PAR (Hollinger et al., 1994; Weiss, 2000; Oliphant et al., 2011). Second, 651 under clear-sky conditions, PAR is mainly direct, resulting in the photosynthesis of sunlit leaves being saturated, 652 whereas the photosynthesis rates of shaded leaves are constrained by the limited intercepted radiation. On cloudy 653 654 days, solar radiation is scattered by clouds in addition to atmospheric aerosols and the proportion of diffuse PAR 655 is high. The saturation effects of the sunlit leaves will be reduced (Gu et al., 2002). Furthermore, photosynthesis and WUE may also benefit from reduced water and heat stress of plants when going from sunny and higher PAR 656 657 conditions to diffuse conditions (Gu et al., 2002; Lloyd et al., 2002; Steiner and Chameides, 2005; Urban et al., 658 2012), especially for water limited ecosystems. Another possible reason for the high photosynthesis rate with 659 diffuse PAR is a change in spectral composition. Diffuse PAR has a higher ratio of blue to red bands than direct PAR, which could stimulate photochemical reactions and stomatal opening (Urban et al., 2012; Cheng et al., 660 661 2015).

662 It has been shown how in temperature-limited systems at high latitudes, incoming longwave radiation (LW_{in}) 663 under cloudy conditions is an important source of energy for snow melting by increasing the land surface

664 temperature (Juszak & Pelliciotti, 2013). We hypothesize that for a given total amount of PAR, increasing diffuse fraction of PAR is associated with increased LW_{in} from clouds and aerosol, which should increase both 665 GPP and ET in temperature-limited conditions. To further explore this, the longwave radiative budget was 666 checked controlling for PAR (Figure 8 (a)). With more clouds, LW_{in} increased significantly (p<0.05) for similar 667 668 incoming PAR. A clear effect of canopy warming (surface temperature, Ts) via longwave radiative budget can 669 be seen in Figure 8 (b). In this temperature-limited ecosystem, increases in surface temperature (e.g. canopy 670 temperature) should enhance photosynthesis. For the air temperature (Ta), we can also see the temperature 671 increase, but the increase of Ta is lower than that of Ts. Moreover, we also check the difference of VPD in 672 diffuse and direct radiation conditions, as shown in Figure 8 (d). It can be seen that VPD in diffuse and direct radiation conditions was not significantly different within similar PAR levels. Even though in the long-term 673 674 dynamics (e.g. path analysis) lower VPD was linked to higher GPP and ET when controlling for PAR, changes in VPD are not significant. Therefore, it can be seen that there is a possible mechanism in this ecosystem that 675 676 increase of long wave incoming due to more clouds could increase the surface temperature and further enhance the photosynthesis rate. This leads to larger GPP and λEc . 677



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Figure 8. Responses of LW_{in} (a), canopy temperature Ts (b) and VPD (c) to various levels of PAR for predominantly diffuse light condition (blue, CI>0.66) and predominantly direct light conditions (red, CI<0.33). The points represent the mean value for specific incoming PAR interval and the error bar represent the significance level at p<0.05). To exclude the effects from phenology and obtain canopy temperature Ts, comparison was conducted with NDVI more than 0.75.

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685 4.5 Comparison between CI and f_{diff}

This study assumes that f_{diff} can be characterized by CI. The reason to use CI is its longer temporal availability 686 687 (2002-2012) at the Soroe flux site, while the observed f_{diff} was available only from 2004. Although previous studies in the Netherlands and in tropical forest ecosystems have proven a strong relationship between f_{diff} and CI 688 (Spitters et al., 1986; Butt et al., 2010), this relationship might be different in high latitude areas. We performed 689 690 a statistical correlation test and model based analysis to check the difference between CI and f_{diff} . Figure 9 (a) 691 shows that CI and f_{diff} are highly correlated (R=0.94). However, there is significant scatter when using the two fractions at the daily time scale. This may be due to that in this study, we assume the PAR is equal to the half of 692 solar shortwave radiation and CI is actually calculated based on the whole range of solar radiation. However, $f_{\rm diff}$ 693 694 is based on the observed diffuse and total PAR on the ground. The effects of scattering are wavelength 695 dependent, therefore CI and f_{diff} may behave differently. Additionally, in the atmosphere, there are two types of 696 wavelength dependent scattering: Rayleigh and Mie scattering. With different types of scattering, the relationship between CI and f_{diff} may be influenced. However, even though the relationship between CI and 697 698 f_{diff} shows some scatter, when we use these two indices as indicators of diffuse radiation in our modeling 699 framework, they show very similar simulation results, as shown in the Taylor Diagram of Figure 9 (b). The simulation results are very similar. We therefore conclude that CI and f_{diff} are very similar and any of these 700 701 quantities could be used to represent the diffuse PAR in this region.



Figure 9. Comparison between f_{diff} and CI from 2004 to 2012. (a) the correlation test between CI and f_{diff} (b) Taylor diagram comparing the GPP and ET modeling performance with CI (X) and f_{diff} (square) and without the diffuse fraction (dot).

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4.6 Other factors that may potentially change responses of GPP and ET to diffuse PAR

This study evaluated the impact of diffuse PAR on GPP and ET, and assessed the interactions between CI and other biophysical environmental variables to jointly regulate GPP and ET. Due to data availability and model complexity, only environmental variables i.e. PAR, Rn-G, LAI, VPD, Ta and SWC were considered in the analysis. Besides these biophysical variables, other variables providing an accurate estimate of the overall fraction of absorbed PAR e.g. leaf inclination angle, leaf optical parameters (reflectance and transmittance) and

leaf-clumping index could influence the impacts of CI on GPP and ET. For instance, Knohl and Baldocchi (2008) 713 found a 20% increase in diffuse PAR effects, when the leaf inclination angle increased from 40 to 70, using a 714 715 canopy radiative transfer model. Variability in the orientation of leave surfaces also changes effects of diffuse PAR. Bonan (2002) suggested that the upper canopy leaves could utilize sunlight more efficiently when they 716 have a near vertical orientation, while the lower foliage must almost be in a horizontal position. Knohl and 717 718 Baldocchi (2008) highlighted that the clumped leave distribution could also have advantages for diffuse PAR. However, our study used top-down and parsimonious GPP and ET models and focused on the ecosystem scale. 719 The impact of leaf properties throughout the whole canopy can thus not be analyzed in detail. The influence of 720 diffuse PAR on evaporation of intercepted rainfall and soil water was also ignored. Diffuse PAR is more 721 722 homogeneous than direct PAR. Diffuse PAR can penetrate deeper and radiation throughout the canopy and at the 723 soil surface is more evenly distributed. This could contribute to not only higher transpiration rate but also more evaporation from soil and the intercepted water. According to Davin and Seneviratne (2012) the response of 724 725 evaporation is less significant than the response of transpiration. This is also supported by our path analysis 726 results. Therefore, in the PT-JPL model, CI was not incorporated in the parameterization of evaporation.

727

728 5. Conclusion

729 The effects of diffuse fraction of PAR on the carbon and water fluxes of a high latitude temperate deciduous forest ecosystem were evaluated using an 11-year (2002-2012) eddy covariance data set from a Danish flux site 730 at Soroe. Using statistical analysis, this study identified that GPP, ET and WUE were mainly controlled by 731 732 variables related with the radiation transfer in the canopy and net energy balance (PAR, LAI and Rn-G) while LUE and EF were primarily controlled by air temperature (Ta). This indicates that this beech forest ecosystem is 733 radiation and temperature limited. Diffuse PAR, expressed by the Cloudiness Index (CI), had positive direct 734 735 effects on GPP, LUE, ET, EF and WUE. In terms of indirect effects, CI mainly interacted with the radiation 736 components in the canopy (PAR, Rn-G and LAI) to influence GPP and ET. Ta and VPD were the major 737 intermediate variables to deliver the indirect influence from CI to LUE and EF. These results indicate that with 738 potentially increasing levels of aerosols and diffuse PAR over the growing season in the future, the WUE, LUE 739 and EF will increase in principle, but the magnitude of the enhancement will depend on the interplay between 740 VPD and Ta.

We tested a joint 'top-down' GPP and ET model, which combines a light use efficiency GPP model (Monteith et 741 al, 1972) and Priestley-Taylor Jet Propulsion Laboratory ET model (Fisher et al., 2008). When incorporating CI 742 into the simulations, the model performance for both GPP and ET improved with the RMSE of the daily GPP 743 decreasing from 1.64 to 1.45 g·C·m⁻²·d⁻¹ (11.68% reduction) and the RMSE of the daily ET decreasing from 744 15.79 to 14.50 W·m⁻² (8.16% reduction). Based on a global sensitivity analysis (GSA), 11.88%, 3.04% and 7.78% 745 746 of the variability of GPP, ET and transpiration, respectively, can be attributed to CI in the growing season from 747 May to October. This proves that CI has largest impacts on GPP, followed by transpiration and finally ET, which results in higher WUE under diffuse fraction conditions. Even though the impact on ET is moderate, it was 748 consistent and we found that most of the ET model improvements when incorporating CI could be linked to the 749 750 transpiration component by comparing with sap flow measurements.

To explain the mechanisms behind GPP and evapotranspiration enhancement with diffuse radiation for fixed levels of PAR, most previous studies have focused on variables affecting the fraction of absorbed PAR. We found that the longwave emission from clouds and aerosols plays an additional role in this high latitude ecosystem. Under diffuse conditions and for same incoming PAR levels, higher longwave emission contributes to higher air and canopy temperature increasing both GPP and transpiration. This highlights the importance of improving the description of the complete radiative transfer in canopies under diffuse and direct conditions in high latitude deciduous forests to model GPP and ET.

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